

One-dimensional inversion of airborne electromagnetic data: application to oil sands exploration

Jamin Cristall, Colin G. Farquharson & Douglas W. Oldenburg

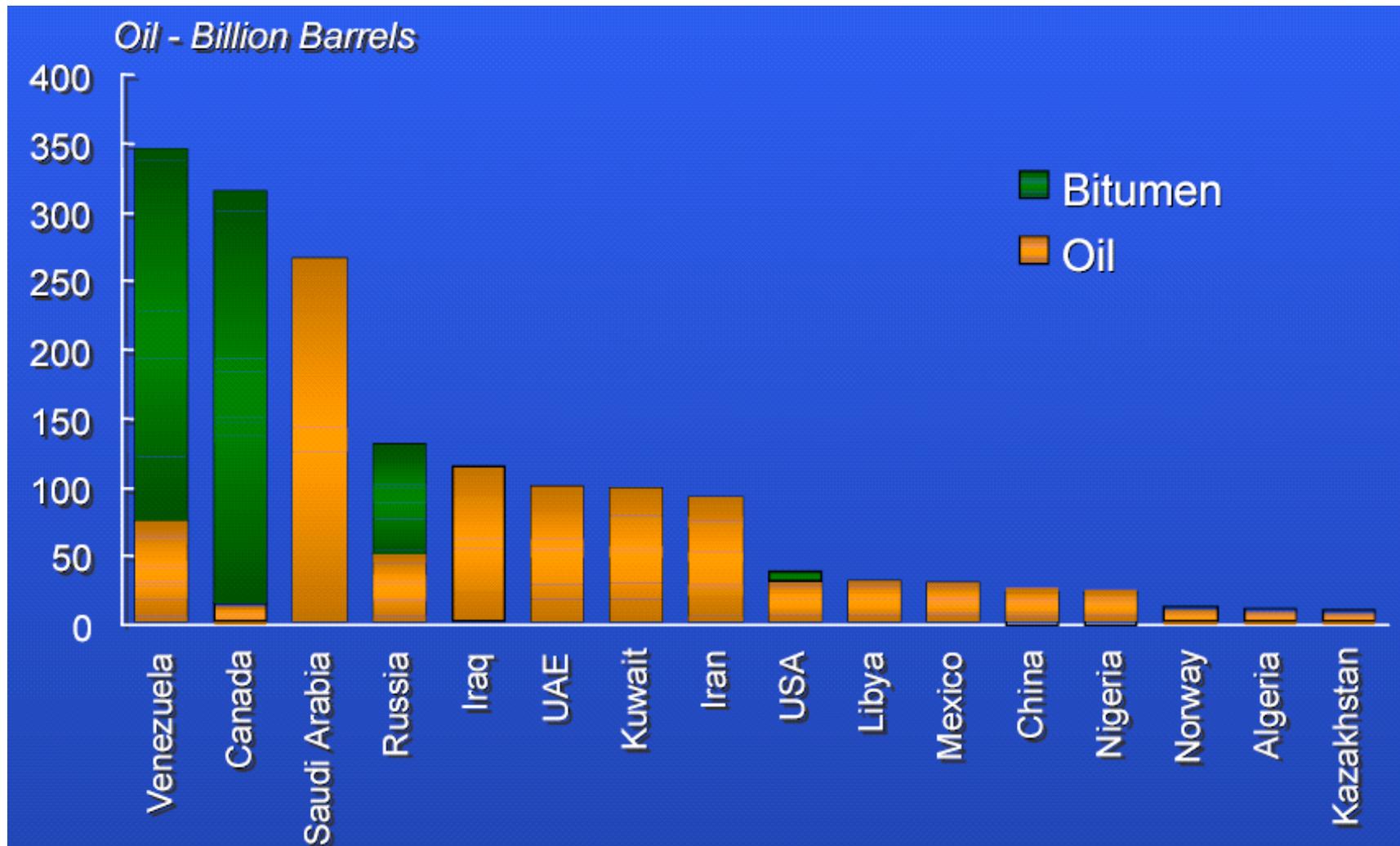
UBC – Geophysical Inversion Facility
Department of Earth & Ocean Sciences
University of British Columbia



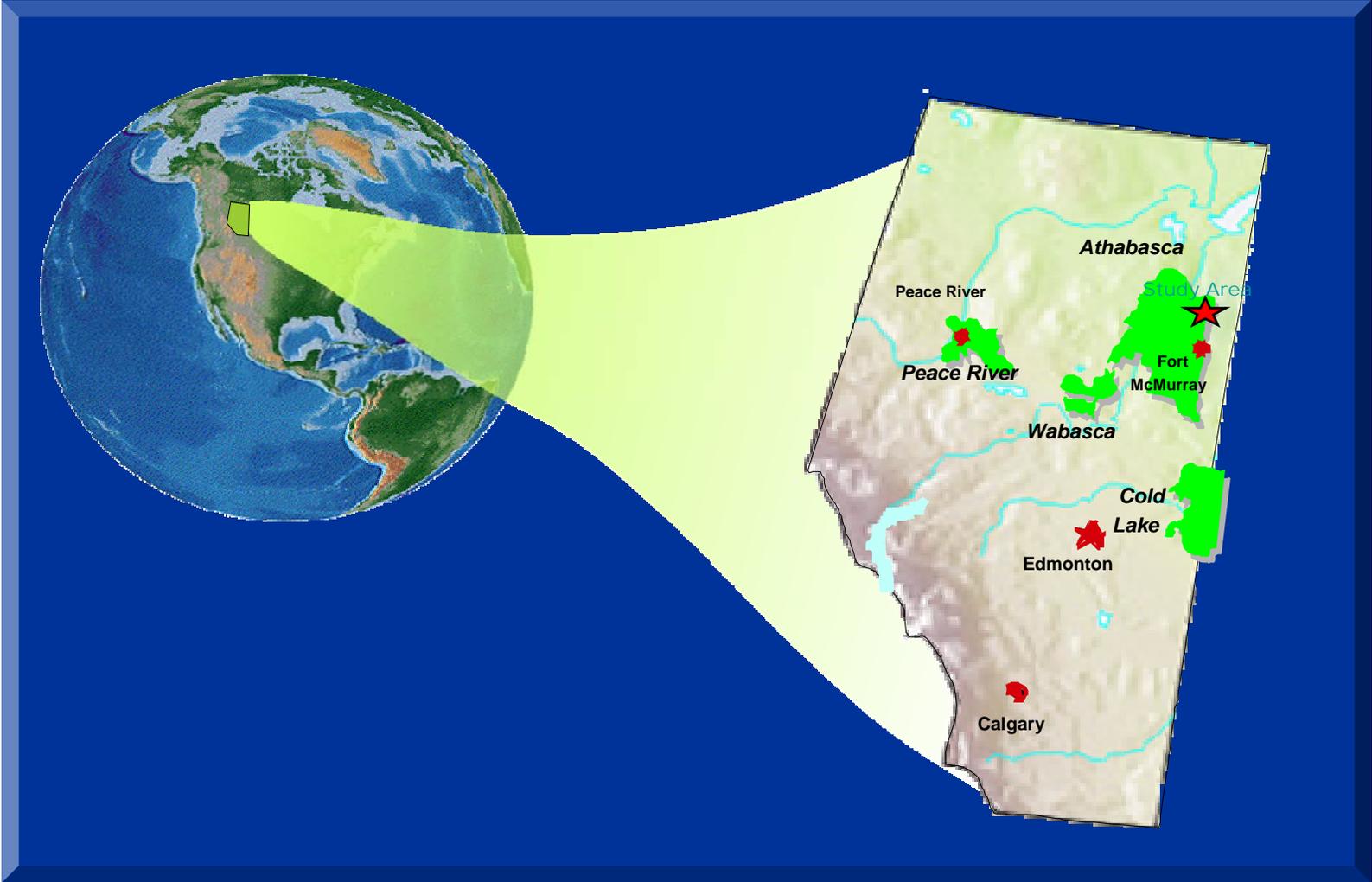
Acknowledgments

- Husky Energy, and Larry Mewhort.
- Richard Kellett, formerly of Komex International.

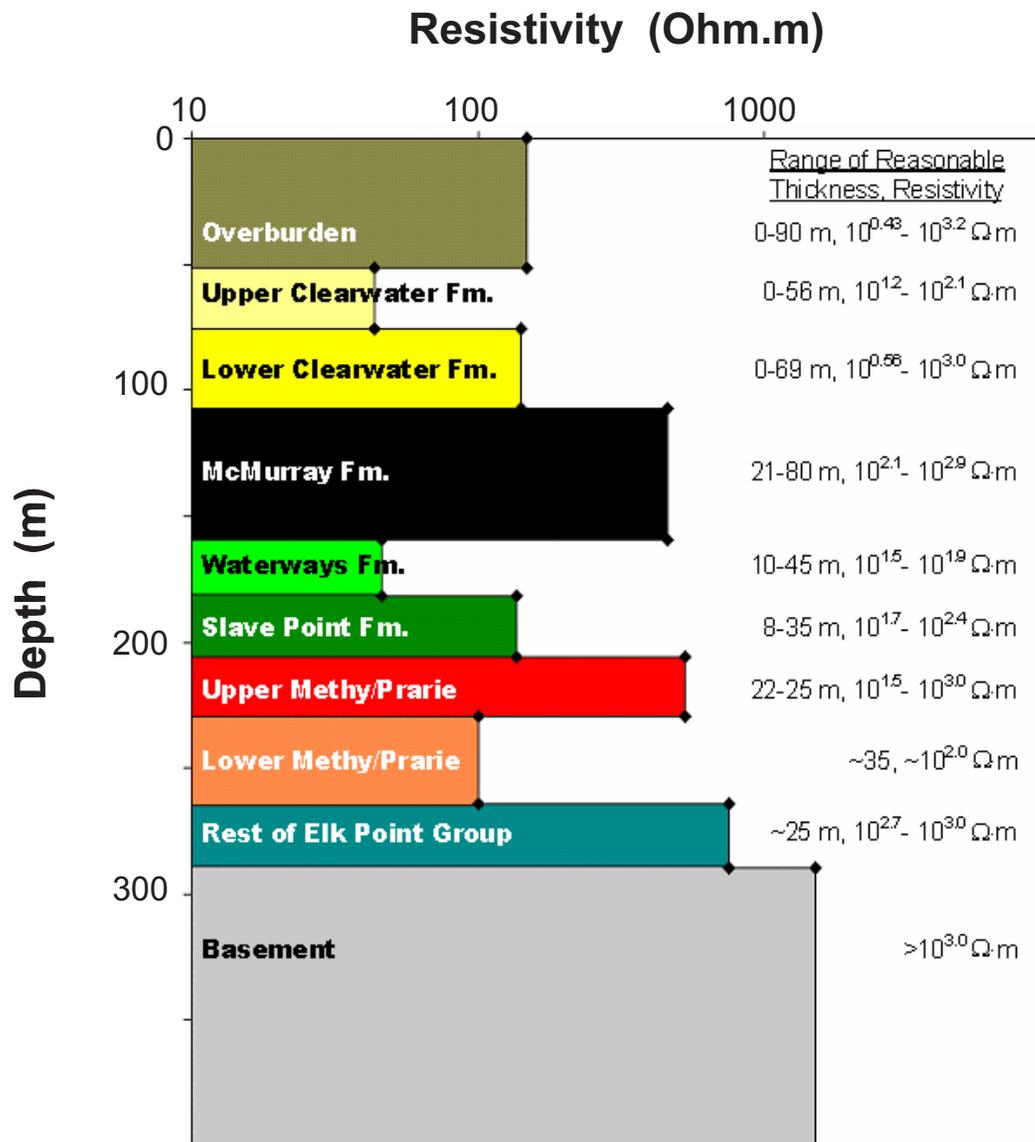
Oil sands in Canada



Oil sands in Canada

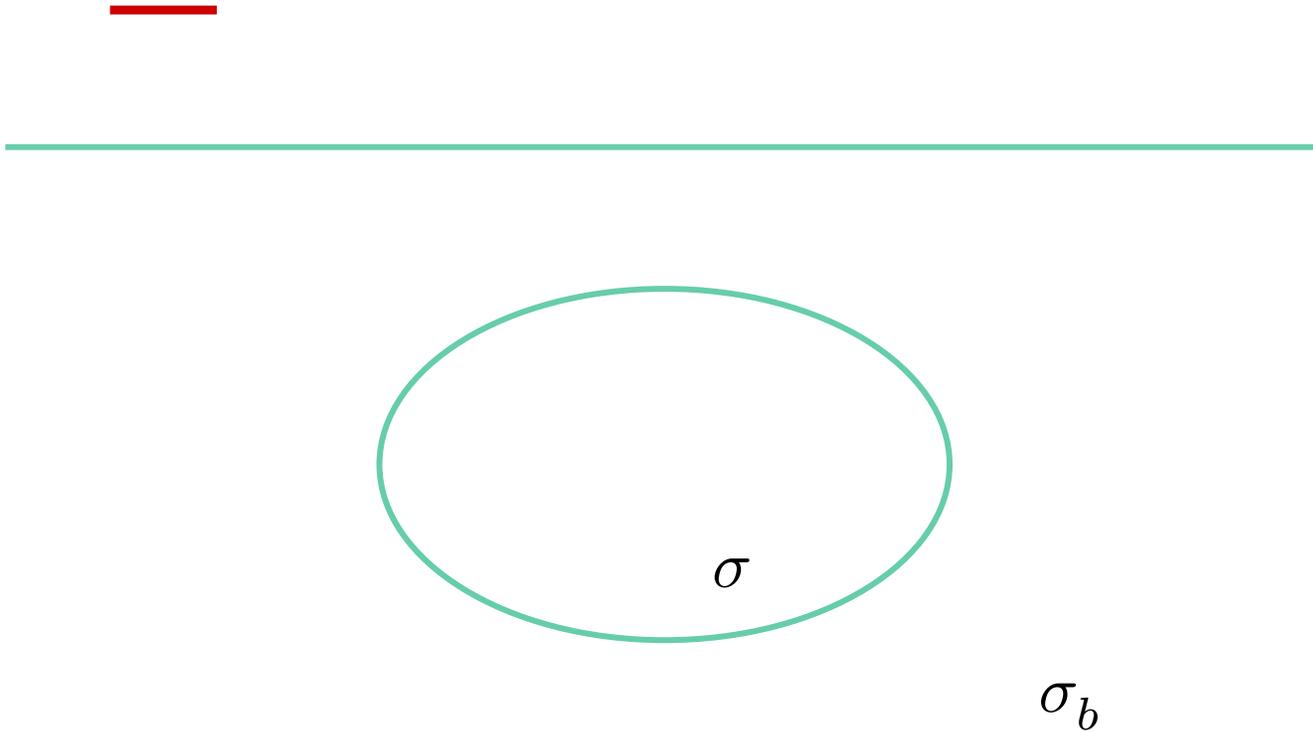


Electrical properties



Electromagnetic induction

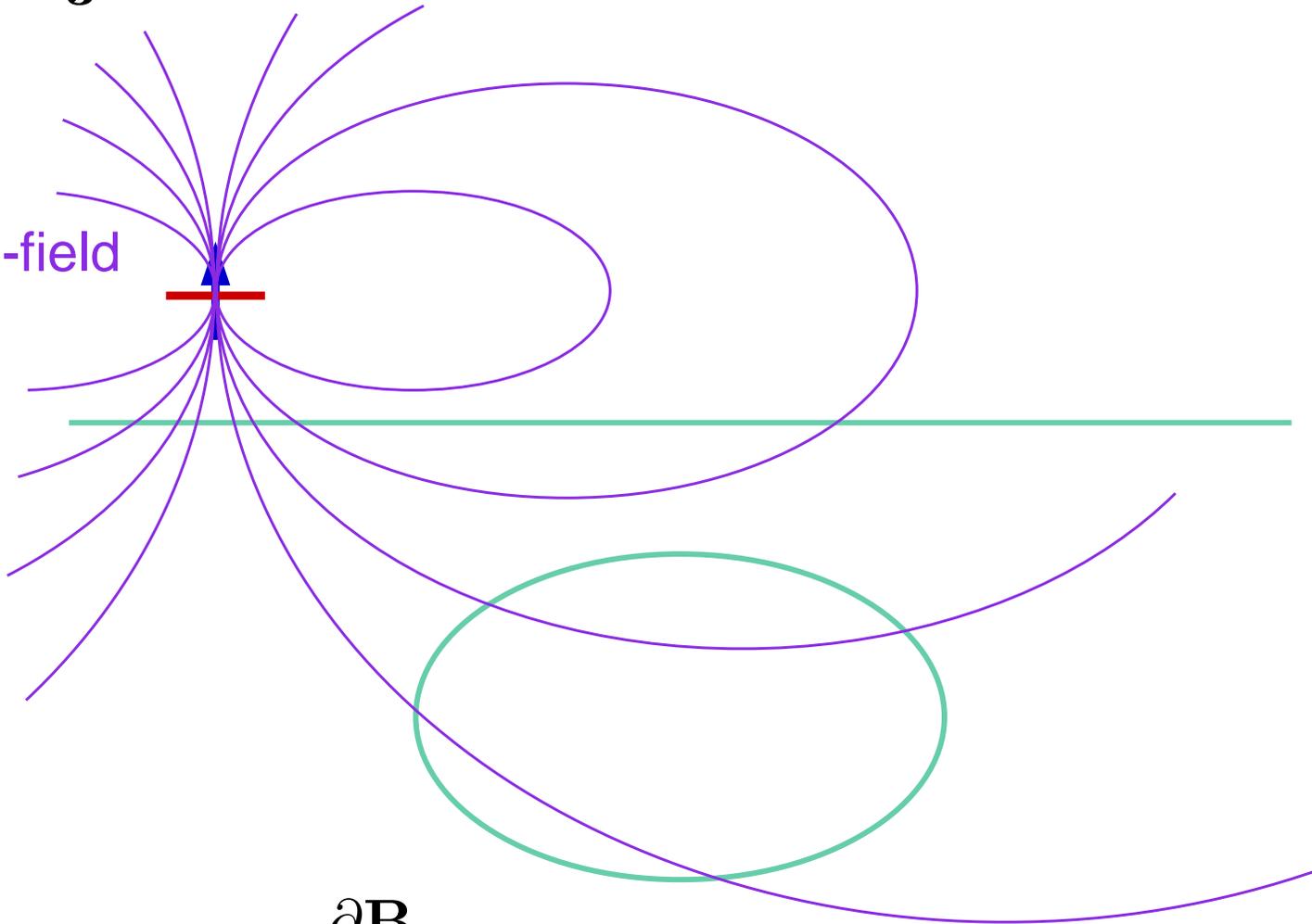
$I(t)$



Electromagnetic induction

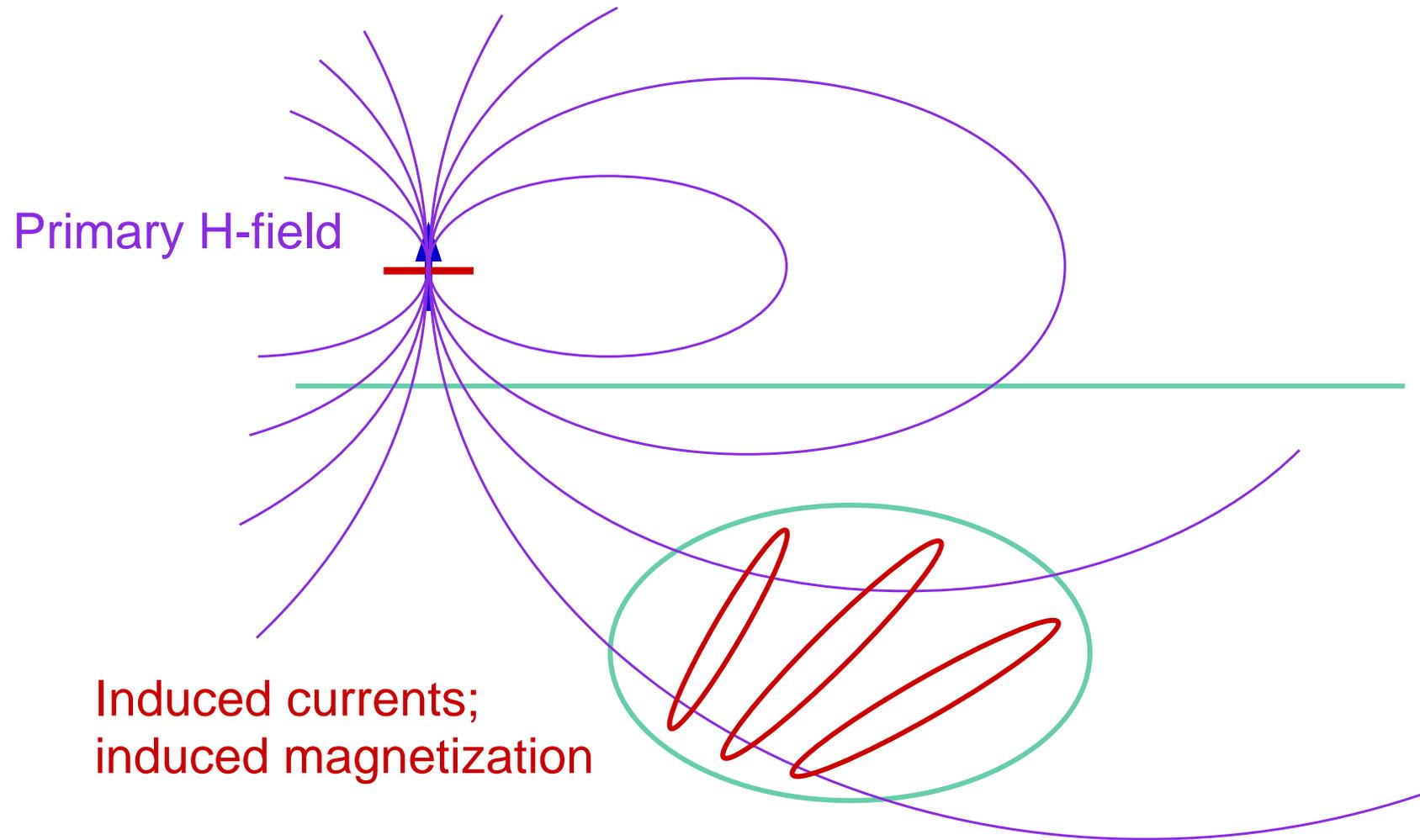
$$\nabla \times \mathbf{H} = \mathbf{J}$$

Primary H-field



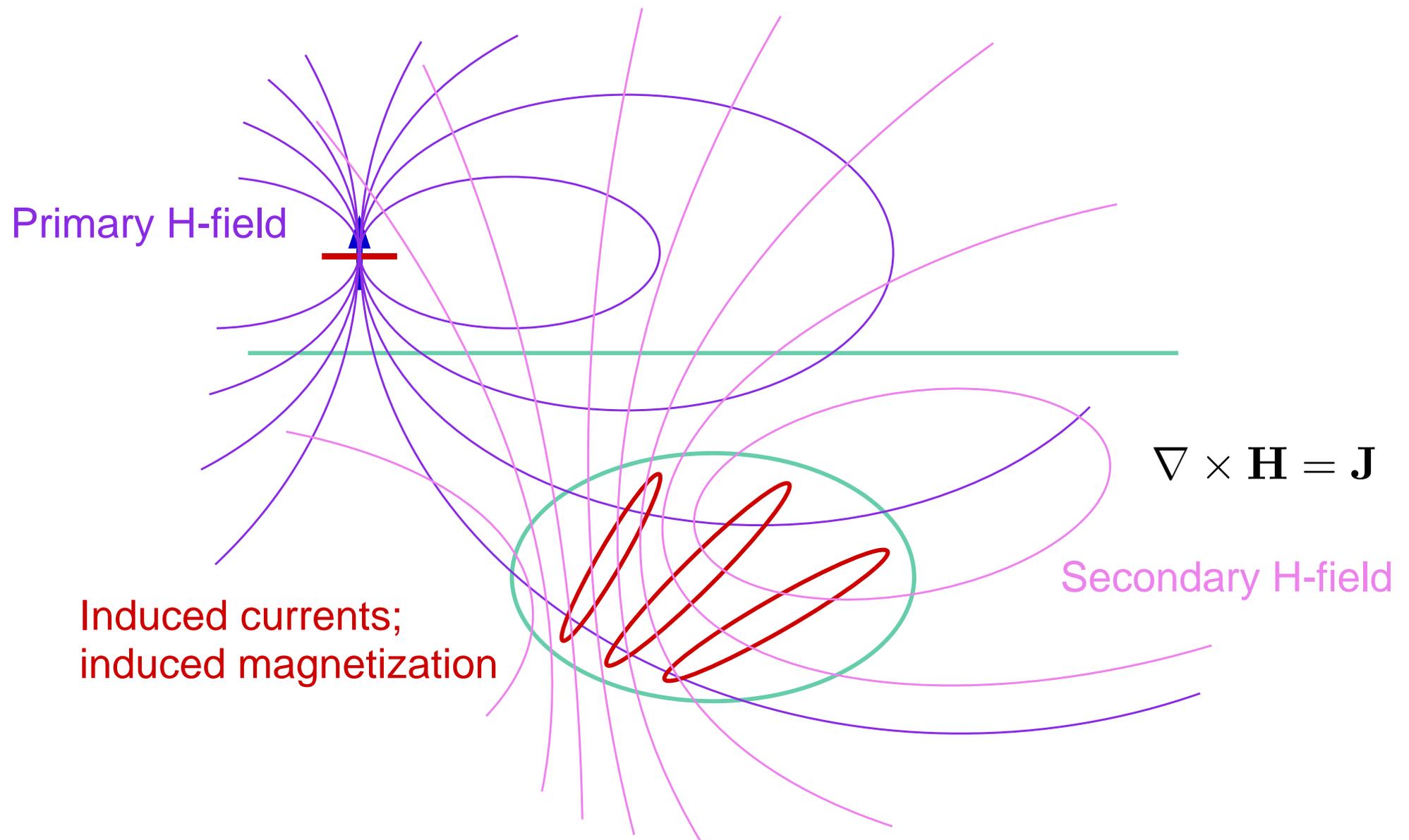
$$\nabla \times \mathbf{E} = -\frac{\partial \mathbf{B}}{\partial t}$$

Electromagnetic induction

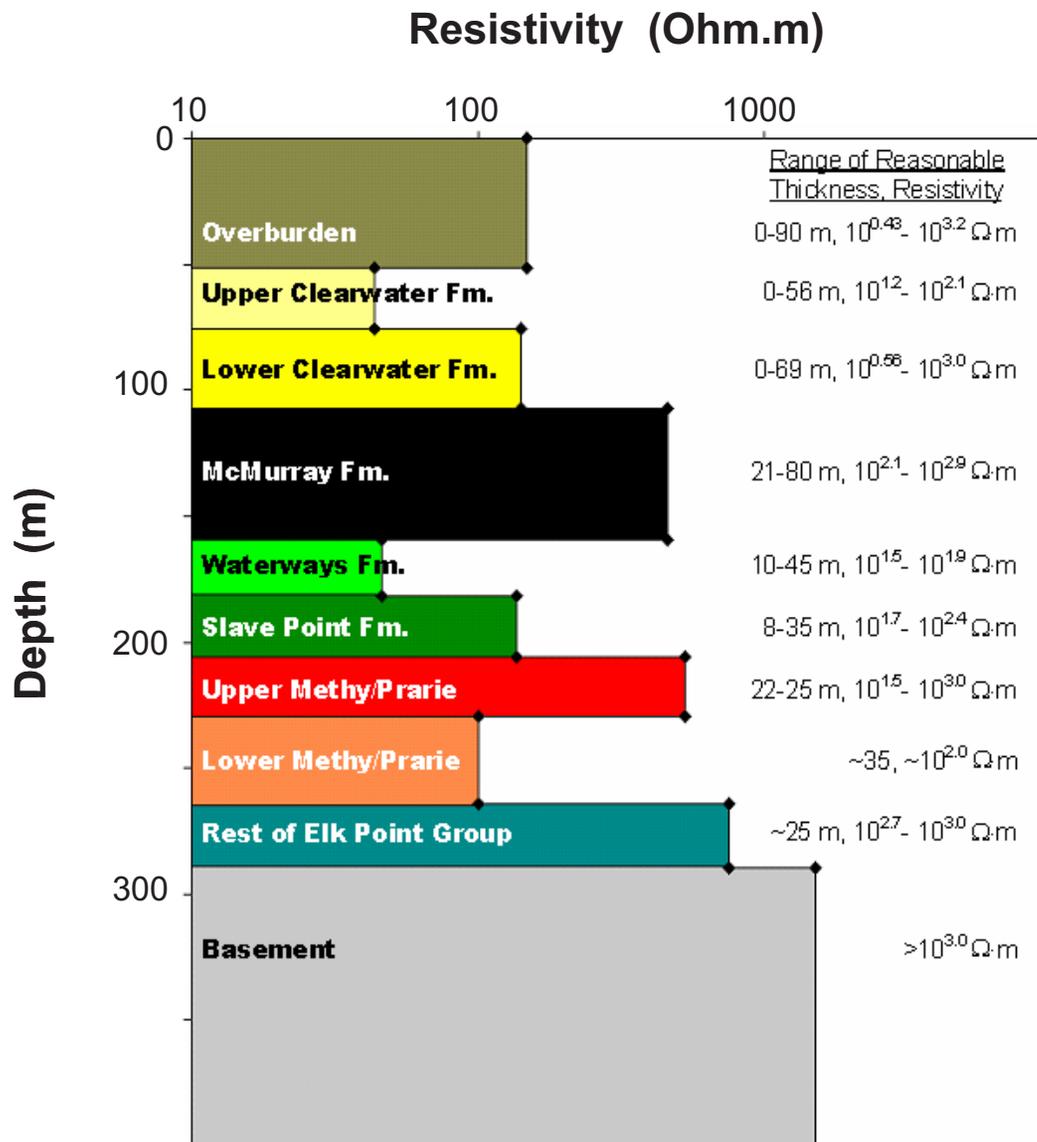


$$\mathbf{J} = \sigma \mathbf{E}; \quad \mathbf{M} = \kappa \mathbf{H}$$

Electromagnetic induction

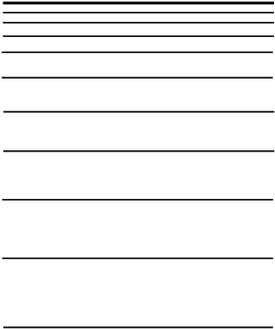


Electrical properties



Inversion methodology: model discretization

- For computations, discretize region of interest:


$$\longleftrightarrow m(\mathbf{r}) = \sum_{j=1}^N m_j \psi_j(\mathbf{r})$$

- * “More linear” inverse problem than when both physical property and cell boundaries are unknowns.
- * To reproduce any spatial distribution, require very fine discretization.
- * Mathematical inverse problem reflects non-uniqueness of original inverse problem.

Inversion methodology: misfit & model structure

- Fitting observations of utmost importance ...

→ measure of misfit: $\phi_d = F\left(\mathbf{W}_d(\mathbf{d}^{\text{obs}} - \mathbf{d}^{\text{prd}})\right)$.

- To deal with non-uniqueness, construct a model with a particular character ...

→ measure of model structure: $\phi_m = F\left(\mathbf{W}(\mathbf{m} - \mathbf{m}^{\text{ref}})\right)$.

Inversion methodology: objective function

- Find the model \mathbf{m} that minimizes the objective function:

$$\Phi = \phi_d + \beta \phi_m.$$

Choose β such that $\phi_d \leq tol$.

- Solution obtained by setting $\nabla_{\mathbf{m}}\Phi = 0$. Hence,

$$\begin{aligned} (\mathbf{J}^T \mathbf{W}_d^T \mathbf{R}_d \mathbf{W}_d \mathbf{J} + \beta \mathbf{W}^T \mathbf{R} \mathbf{W}) \delta \mathbf{m} = \\ - \mathbf{J}^T \mathbf{W}_d^T \mathbf{R}_d \mathbf{W}_d (\mathbf{d}^{\text{obs}} - \mathbf{d}^n) - \beta \mathbf{W}^T \mathbf{R} \mathbf{W} (\mathbf{m}^n - \mathbf{m}^{\text{ref}}), \end{aligned}$$

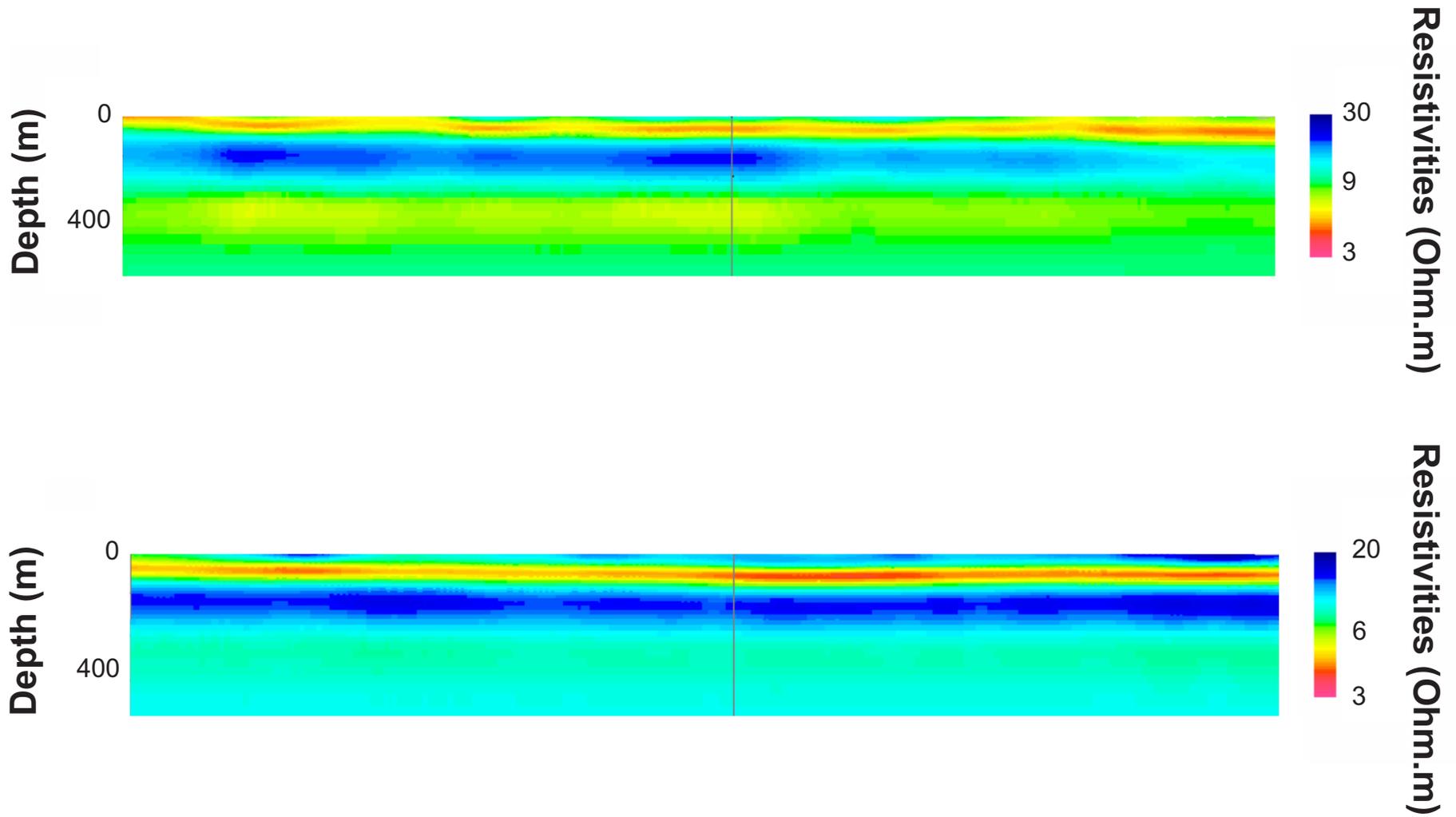
where \mathbf{J} is the Jacobian matrix of sensitivities:

$$J_{ij} = \frac{\partial d_i}{\partial m_j}.$$

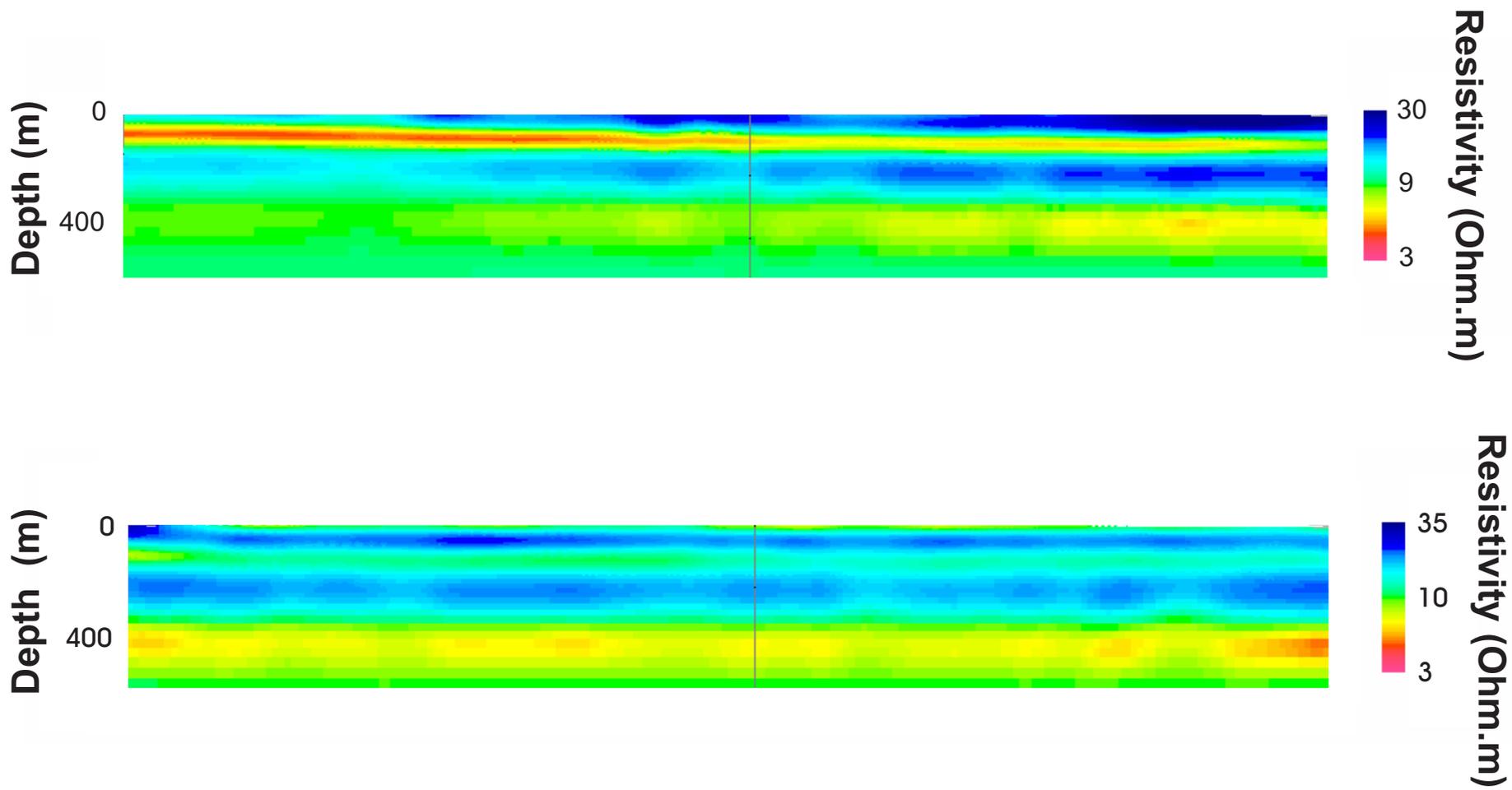
GEOTEM data from the Fort McMurray area



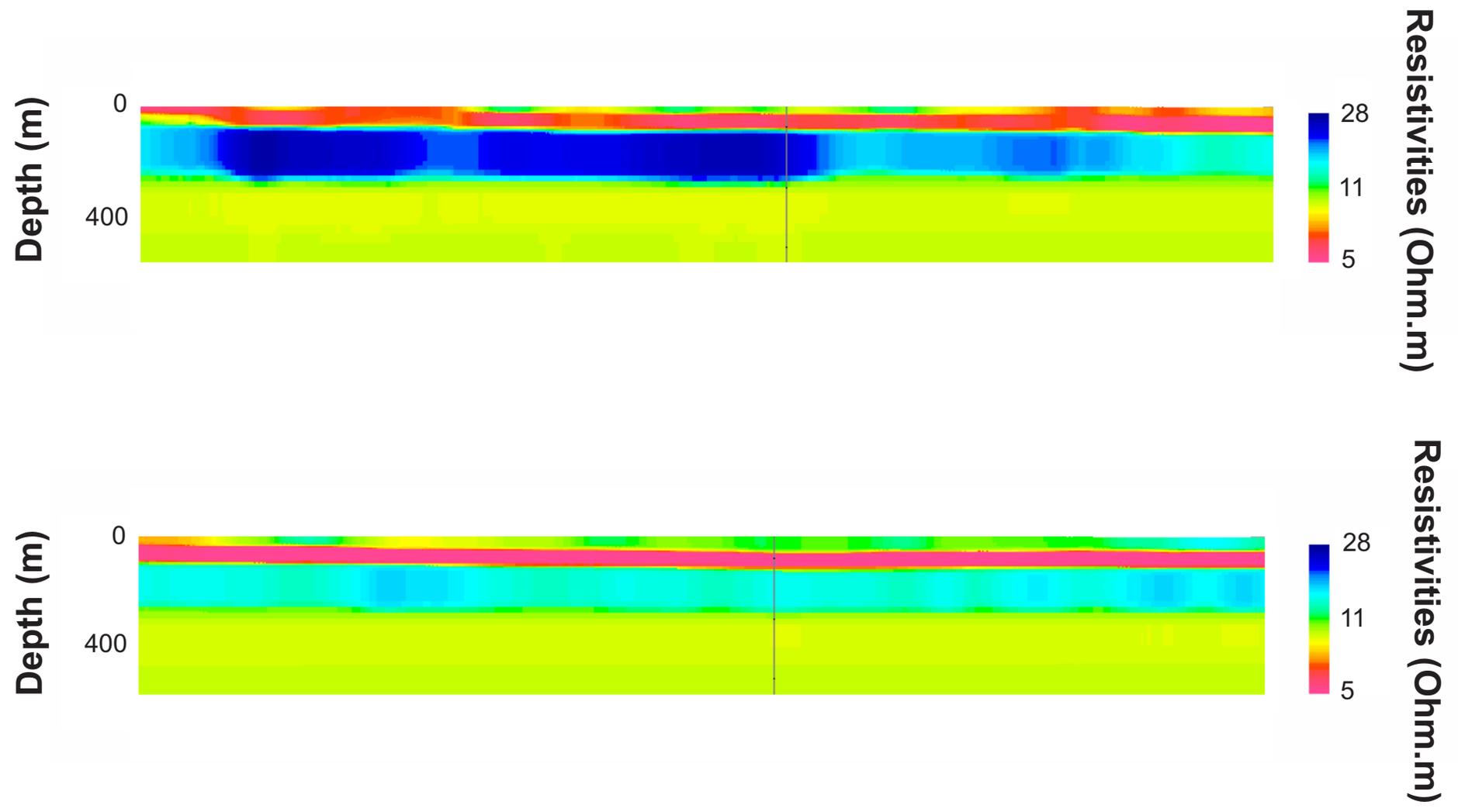
Inversion results – line 10101 – 12



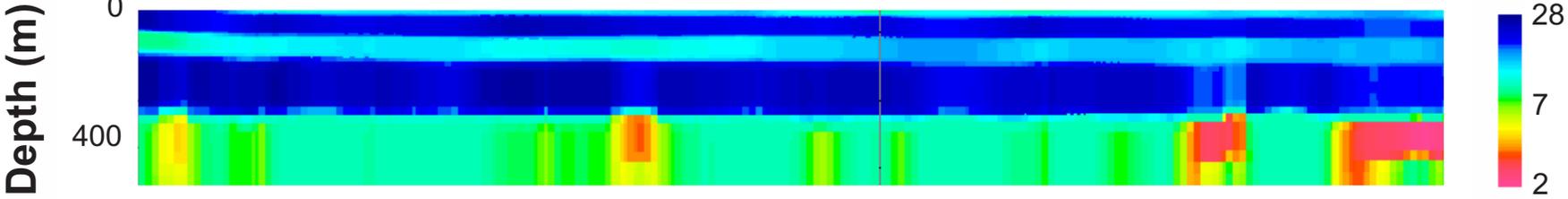
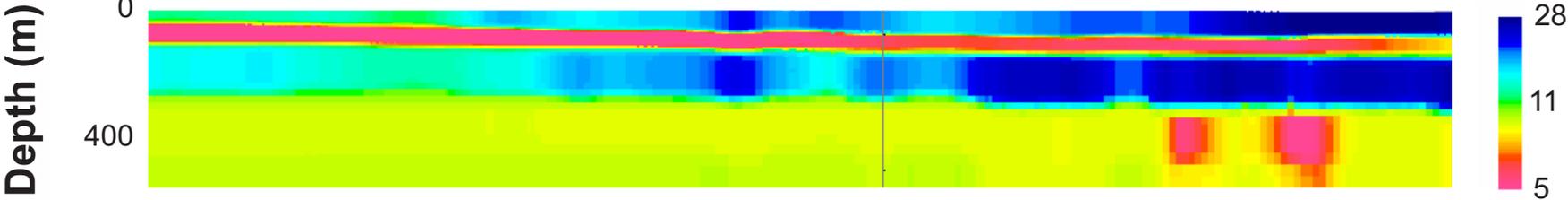
Inversion results – line 10101 (contd) – 12



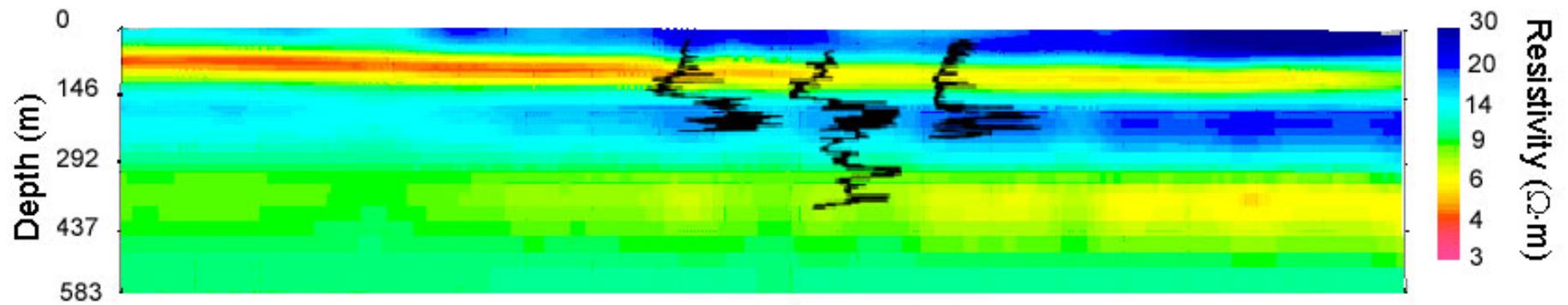
Inversion results – line 10101 – 11



Inversion results – line 10101 (contd) – 11



Inversion results – comparison with well-logs



Summary

- ★ A new, atypical application of electromagnetics in the investigation of the Earth's subsurface.
- ★ Quantitative interpretation via inversion (even 1D).